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Estimation of drought trends and comparison between SPI and SPEI with prediction using machine learning models in Rangpur, Bangladesh

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ABSTRACT

This study investigates drought trends, SPI-SPEI comparisons, and predictions in Rangpur, Bangladesh, from 1979 to 2020. We employed Modified Mann-Kendall for trend analysis, SPI and SPEI for drought assessment, and Pearson Correlation Coefficient and Simple Linear Regression for evaluating SPI and SPEI relationships. Additionally, we utilized ANN, SVM, and RF for prediction. The study revealed notable negative trends in seasonal and annual drought, with the highest z statistics observed for SPI 06 (-2.75), SPI 09 (-4.50), SPI 12 (5.60), SPI 24 (-8.40), SPEI 06 (-5.13), SPEI 09 (-6.82), SPEI 12 (-8.04), and SPEI 24 (-11.20). Strong correlations were identified across all SPI and SPEI indices, with coefficients peaking at 97%, 98%, 98%, and 97% for 06, 09, 12, and 24-month periods, respectively. The comparative assessment favored SPEI over SPI, highlighting its superiority and accuracy. The ANN prediction model showed significant results for short-term and seasonal drought forecasts, projecting SPEI 03 and SPEI 06 increases of 0.02 and 0.24, respectively. However, long-term drought estimation exhibited insignificant performance across all predictive models. This emphasizes the need for developing essential predictive tools for future drought variability.

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1. Introduction

The onset and duration of drought vary widely depending on the region, and is influenced by various climatic and atmospheric factors (Sheffield & Wood, 2008). In addition to a prolonged absence of rainfall, drought conditions are often associated with elevated temperatures, strong winds, and low moisture levels, which can profoundly impact both the natural and human systems (Mekonnen & Gokcekus, 2020). The effects of drought are felt globally, with an estimated 55 million people being impacted each year, and the most significant threat posed to crops, animals, and human populations in virtually every region (Shahfahad et al., 2022). The consequences of drought can be severe, including loss of livelihoods, increased risk of illness and death, and mass migration (Zinat et al., 2020). Thus, understanding the causes and impacts of drought is crucial for developing effective strategies to mitigate its effects and ensure the resilience of vulnerable communities.

The global water crisis poses a significant threat to the livelihoods and well-being of up to 700 million people, or 40% of the world's population (Islam et al.,

2017). By 2030, millions may be forced to leave their homes due to severe drought conditions (Islam et al., 2017). Bangladesh is one of the most vulnerable countries in the world, facing a range of socioeconomic, geographic, and climatic challenges that leave it less resilient to the impacts of natural disasters (Kamruzzaman et al., 2022; Uddin et al., 2020). Despite its relatively high precipitation levels, the country experiences a dry period each year that can have severe consequences for local communities and the economy (Alam et al., 2021). The northern part of Bangladesh, with its diverse range of precipitation patterns and high concentrations of people, is particularly at risk of experiencing the effects of drought (Banglapedia, 2014).

Drought has been recognized as an ecological emergency and has been the subject of intense study across multiple scientific disciplines, including natural sciences, ecology, hydrology, meteorology, geology, and agriculture (Mainuddin et al., 2020; Mortuza et al., 2019; Rahman & Azim, 2021, 2022; Rahman et al., 2021). The effects of drought can be far-reaching and devastating, exacerbating job insecurity, reducing access to food, and causing abnormal price

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spikes for rural residents. Extreme drought can result in crop losses, food shortages, and the loss of livestock and biodiversity, leading to migration in affected areas (Rahman & Azim, 2021). Therefore, it is crucial to better understand the present and future drought scenarios in Bangladesh and adopt effective management strategies to mitigate its impacts. Bangladesh is among the countries that are most vulnerable to weather-related disasters. Climate change is exacerbating the effects of droughts and monsoons in Southeast Asia, potentially leading to increased intensity of both flooding and dryness in the region. The National Drought Mitigation Center reported that the frequency of droughts in Bangladesh has increased over the past several decades. The northern part of the country is particularly susceptible to prolonged droughts due to its high variability in rainfall patterns (von Neumann, 1941). The prolonged drought in the northern region presents significant environmental, social, and economic challenges and highlights the need for effective strategies to mitigate the impacts of drought (Alexandersson, 1986).

Droughts have significantly impacted various aspects of life in Bangladesh, including agriculture, food production, land degradation, economics, and communities (Jerin et al., 2021). In particular, drought stress during crop growth can significantly affect crop production in Bangladesh and other LDCs in Asia (Uddin et al., 2020). Several studies have been conducted to assess and analyze the risks of drought in Bangladesh. The study by Rafiuddin (von Neumann, 1941) used the Standardized Precipitation Index (SPI) to detect dry spells and concluded that the use of SPI throughout an area provides better consistency of drought conditions compared to using single station data. The 1994 drought, which lasted from the end of 1994 to the start of the monsoons in 1995, was the worst in a century and had severe impacts on the country's northwestern region (Kendall, 1975). The effects of drought on the ecosystem, human health, society and economy, and agriculture have been explored in previous research (Hamed & Rao, 1998; McKee et al., 1993; Rahman & ARMT, 2019; Rahman & Azim, 2021; Wang et al., 2019; Zannat et al., 2019). For example, Alam (Pramudya & Onishi, 2018) evaluated the occurrence of dryness in northern Bangladesh and its relationship with water stress parameters and population health in a vulnerable region. Islam et al. studied northern and southern drought conditions using SPEI where no comparisons and predictions between SPI and SPEI is employed (Islam et al., 2017). Only two different studies are found for comparing SPI and SPEI where (Uddin et al., 2020) focused only three zones and ignored micro level study (Kamruzzaman et al., 2022); investigated different zones using traditional Mann-Kendall test for drought trend. Therefore, the explained

literature investigates the effect of dry spells, as evaluated by such SPI, on daily death rates from natural causes, circulatory-related causes, and respiratory-related causes in northern Bangladesh between 2007 and 2017 (Alam et al., 2021). However, the northern region, particularly in Rangpur, has never been the focus of systematic drought research utilizing SPI and SPEI while revealing drought trends using the Modified Mann-Kendall test (MMK). In addition, no study contributed to the working efficiency of SPI and SPEI with their future variability. Hence, this study evaluated and predicted drought patterns in the northern regions of Rangpur in Bangladesh. The objectives of the study are:

- (1) To identify temporal drought trends in the Rangpur district during 1979-2020;
- (2) To determine the relationship and comparative assessment between the SPI and SPEI in drought variability;
- (3) To predict temporal drought in the Rangpur district.

The sections of the study include firstly revealing drought trends using MMK, then finding the drought conditions and employing comparisons using SPI and SPEI, and finally predicting the drought variability.

2. Data and methods

2.1 Study area description

The northern region's Rangpur district was chosen as the study location due largely to the availability of specific climatic zone there (Figure 1). The area of the Rangpur district is 2,641.84 sq. km. Nilphamari and Gaibandha districts lie north, Kurigram and Dinajpur to the east, and Dinajpur to the west of this area. Weather-wise, Bangladesh has four different seasons: winter (December-February), pre-monsoon (March-May), monsoon (June-September), and post-monsoon (September-December) (October to November) (Banglapedia, 2014). Most of the country's territories have humid to subhumid climates with anomalous spatial and temporal patterns. The Rangpur district's digital elevation model (DEM) is also shown in Figure 1; the area is mostly plain, with some elevated regions areas, the province's northern and northeastern corners. The mean annual precipitation is 2428 mm, and more than 80% of precipitation falls during the monsoon season (Mainuddin et al., 2020). The Bangladesh Meteorological Department (BMD) stated that the yearly average temperature was 25.7°C. In addition, the highest temperatures registered during the summer vary from 30°C to 40°C, and the minimum temperatures during the winter vary from 5.2°C to 10°C, notably in January (Rahman & Azim, 2022). The weather patterns in

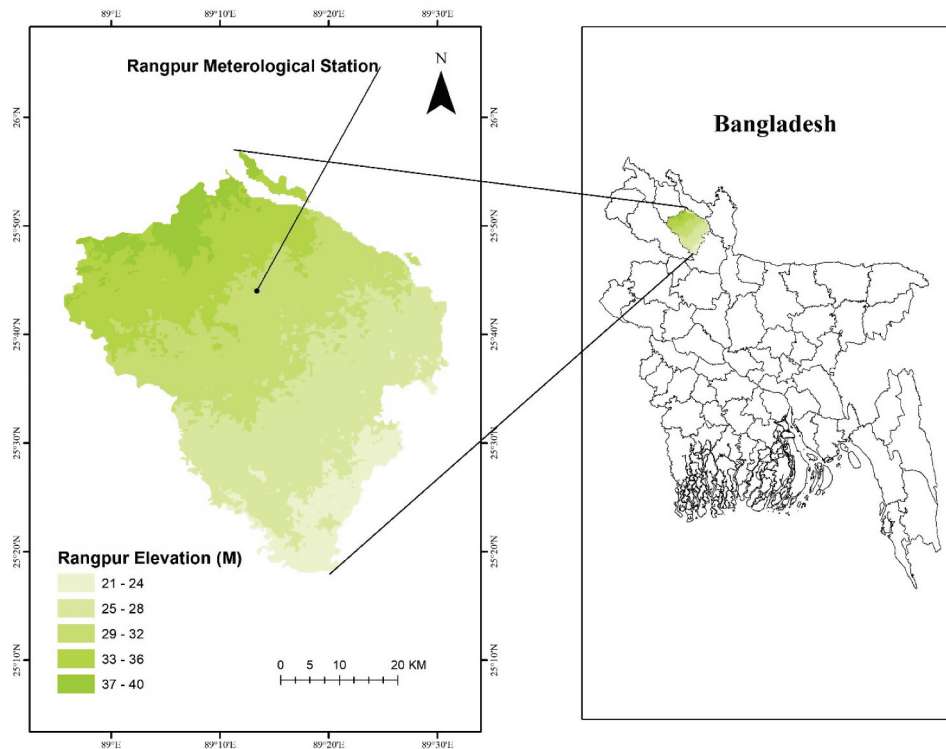


Figure 1. Study area of Rangpur district (Digital elevation model of Rangpur, Bangladesh). **Source:** Modified from humdata.org

Rangpur is thus classed as warm and temperate. In Rangpur, the summertime is significantly rainier than the winter. This climate has been classed as Cwa by Koppen and Geiger. The mean temperature in Rangpur is 24.3°C, and the annual precipitation is 2498 mm or 98.3 inches (Mainuddin et al., 2020).

2.2 Data sources and quality control

The monthly recorded rainfall observations were obtained from Bangladesh Agricultural Research Council (BARC) and Bangladesh Meteorological Department (BMD) for the years 1979 to 2020 (42 years). Additionally, from 1979 to 2018 data was acquired from the freely available source of BARC and from 2019 to 2020, data were collected from BMD. Moreover, only one station covers the northern part of the northern climatic zone and the entire Rangpur (Rahman & Azim, 2021). Buishand and von Neumann introduced the significance of testing homogeneity for observation data (Buishand, 1982; von Neumann, 1941). Additionally, datasets checked for homogeneity and homogenous (Alexandersson, 1986; Buishand, 1982; von Neumann, 1941). When utilizing historical information, missing data is a crucial concern. Two percent of the missing value is widely accepted because of its inefficiency in influencing outcomes where we found < 1% missing value (Mortuza et al., 2019; Rahman et al., 2021).

2.3 Methods

We performed linear regression as a part of the parametric test to detect the trend of monthly rainfall before applying non-parametric tests, including Modified Mann-Kendall for trend analysis, SPEI, and SPI for drought as well as dry and wet period variability (Hamed & Rao, 1998; Jerin et al., 2021; Kendall, 1975). Moreover, Pearson's Correlation Test was employed to illustrate a link between SPEI and SPI (Wang et al., 2019). To predict the future variability of drought tendency patterns, we utilized the training and testing of these prediction methods were carried out using 70% and 30% of the data, respectively. The subsequent section of this article briefly presents the statistical phenomenon.

2.4 Modified Mann Kendall Trend test (MMK)

Hamed & Rao (1998) proposed the Modified Mann-Kendall (MMK) test, which considers autocorrelation in time series data when assessing trends. The variance of the MK statistic ($Var(S)$) is modified accordingly (Rahman & ARMT, 2019; Zannat et al., 2019). The autocorrelation of the data is calculated using the following equations:

$$Var(S) = Var(S) * \left(\frac{n}{n^*}\right)$$

$$\left(\frac{n}{n^*}\right) = 1 + \left(\frac{2}{n(n-1)(n-2)}\right) * \sum_{k=1}^{n-1} (n-k)(n-k-1)(n-k-2)^r .k$$

$$r_k = \frac{\left(\frac{1}{n-k}\right) \sum_{i=1}^{n-k} (x_i - x)(x_{i+k} - x)}{\left(\frac{1}{n}\right) \sum_{i=1}^n (x_i - x)^2}$$

Here, $Var(S)$ is calculated via $Var(S)^*$, nn^* represents the adjusted equation of auto-correlated statistics, rk is the auto-correlation coefficients of the k th lag, and x defines the temporal-sequence records mean. Using the following equation, the relevance of the autocorrelation parameter of the k th lag at a 95% confidence interval is calculated:

$$\left(\frac{-1 - 1.96\sqrt{n-k-1}}{n-k}\right) \leq r_k(95\%) \leq \left(\frac{-1 + 1.96\sqrt{n-k-1}}{n-k}\right)$$

Whereas if acquired rk matches the aforementioned scenario, data with a 95% level of certainty may be retrieved. In order to discern the pattern in the response variable, it is imperative to eliminate the impact of auto-correlation across various temporal delays, provided that the variables are dependent. Moreover, the tendency in time series data can be inferred if Z falls below or exceeds the Z value of the normal distribution at a 95% level of significance.

2.5 Standardized Precipitation Index (SPI)

In drought assessment, the Standardized Precipitation Index (SPI) is a widely utilized method that estimates rainfall based on various timelines, including one, three, six, nine, twelve, and twenty-four months (McKee et al., 1993; Pramudya & Onishi, 2018; Shahid, 2008). This approach involves comparing time series rainfall records for a specific period to sums from a corresponding period, thus facilitating the analysis of tropical and temperate events. To calculate the SPI, time series rainfall records from the Rangpur weather station were fitted to a gamma probability distribution function:

$$\gamma(x) = \frac{1}{\beta^\alpha \Gamma(\alpha)} x^{\alpha-1} e^{-\frac{x}{\beta}} \text{ for } x > 0$$

When the factor is more significant than zero, the curve is also greater than zero, indicating a positive level of rainfall. The gamma feature can be specified by alpha. However, the likelihood formulae used do not effectively calculate the gamma unit parameters

Table 1. Drought index based on SPI values.

Category	SPI	Probability of Occurrence (%)
Extreme drought	-2.00 and less	2.3
Severe drought	-1.50 to -1.99	4.4
Moderate drought	-1.00 to -1.49	9.2
Near normal	.99 to -.99	34.1
Moderate wet	1.00 to 1.49	9.2
Severely wet	1.50 to 1.99	4.4
Extremely wet	2.0 and above	2.3

Table 2. Drought index based on SPEI values.

Category	SPEI Extremes
Extreme drought	-1.83 and less
Severe drought	-1.82 to -1.43
Moderate drought	-1.42 to -1.49
Near normal	-1.0 to 1.0
Moderate wet	1.42 to < 1.0
Severely wet	1.82 to < 1.43
Extremely wet	1.83 and above

alpha and alpha for each location and time period more significant $\alpha = \frac{1}{4A} \left(1a + \sqrt{1 + \frac{4A}{3}}\right)$

$$\beta = \frac{\bar{x}}{\alpha}$$

where $A = \ln(\bar{x}) - [\sum \ln(x)/n]$, n = quantity reflects measurements of rainfall. As a result, the amount of precipitation there at stations may be accurately characterized by a function of computational incremental probabilities, as shown below:

$$G(x) = \int_0^x g(x) dx = \frac{1}{\beta^\alpha \Gamma(\alpha)} \int_0^x x^{\alpha-1} e^{-x/\beta} dx$$

Because the simulated data is confusing for $x = 0$ and a rainfall spectrum may contain zeros, the accumulated probability is: $H(x) = q + (1 - q)G(x)$. Where q is the probability of zero. This SPI is calculated by converting its continuous likelihood $H(x)$ towards the probability of the ordinary distribution (McKee et al., 1993). It enables the calculation of an identifier property for SPI. McKee (Zou et al., 2003) classified wet and dry instances that matched SPI characteristics, as shown in Table 1.

2.6 Standardized Precipitation Evapotranspiration Index (SPEI)

The Standardized Precipitation Evapotranspiration Index (SPEI) is a relatively new meteorological drought measure that combines the benefits of the Standardized Precipitation Index (SPI) with temperature variations. It has been suggested that SPEI is more suitable for drought monitoring and assessment in the context of climate change (Mohsenipour et al., 2018; Vicente-Serrano, 2015). Moreover, SPI lacks hydrological balance and has been used only to adjust statistical regions to specific objectives (WANG et al.,

2018). To address these empirical shortcomings, SPEI is utilized as it incorporates the sensitivity of the Palmer Drought Severity Index (PDSI) to fluctuations in evaporative demand caused by temperature variations and trends with simpler SPI simulations (Vicente-Serrano, 2006). To estimate the SPEI, one could firstly determine the periodic potential evapotranspiration (PET), and after that identify the disparity in the periodic water level of a particular month (i.e., deficiency D_i) by deducting PET out from rainfall measurement of that month (P_i).

$$D_i = P_i - PET_i$$

Once a log-logistic likelihood functional model is applied to the normalized deficiency response variable, the SPEI level may be computed as the confidence interval of the accumulated likelihood of D_i .

$$SPEI_i = w_i - \frac{2.515517 + 0.802853W_i + 0.010328W_i^2}{1 + 1.432788W_i + 0.189269W_i^2 + 0.001308W_i^3}$$

$$w_i = \sqrt{-21np} \text{ for } p < 0.05 \text{ and}$$

$$w_i = \sqrt{-21n(1-p)} \text{ for } p < 0.05$$

When p is greater than 0.5, the symbol of such resulting SPEI flips.

The different time scales of SPI and SPEI have been employed, including 3-months, 6-months, 9-months, 12-months, and 24-months period (V-S et al., 2010). The 3-month period is a useful tool to assess short-term drought trends over a three-month period (Rahman et al., 2021; Vicente-Serrano & Beguería, 2016). However, it is crucial to examine the data over a longer time frame to avoid any misinterpretations about droughts, as a relatively normal 3-month period may occur within a more extensive drought (Rahman et al., 2021). Additionally, a 3-month period may be significantly useful for medium or seasonal drought pattern extraction which may help in cropping activities (Rahman et al., 2021). Furthermore, long-term droughts are identified by 9, 12 and 24 months period that can aid agricultural productions, and drought mitigation strategy development (Islam et al., 2017; Rahman et al., 2021; Shahfahad et al., 2022).

2.7 Pearson correlation coefficient

This study employs the Pearson correlation coefficient statistical technique to examine the correlation between SPI and SPEI (Tefera et al., 2019). Excel to calculate the correlation coefficient, considering the relevant parameters that constitute the Pearson mechanism. To perform a Pearson correlation test, the first array must contain distinct numerical quantities, while the second array should consist of variables to be analyzed. Only numeric or string values, array parameters, or pointers to integers used as

responses are acceptable. Units with the quantity 0 can be retained, even if they contain text, resulting in either real or null strings, whereas units with other attributes are discarded. If both arrays are null or contain a variety of observations, the PEARSON function will generate the #N/A estimation error. The Pearson correlation coefficient, r , can be calculated as follows:

$$r = \frac{\sum (x - \bar{x})(y - \bar{y})}{\sqrt{\sum (x - \bar{x})^2 \sum (y - \bar{y})^2}}$$

where x and y indicate first and second array, respectively.

2.8 Simple linear regression

In this research article, we utilized Simple Linear Regression to conduct a comparative analysis of SPEI and SPI (Zou et al., 2003). The Simple Linear Regression equation that was employed in this study is as follows:

$$y = bx + a + \varepsilon$$

The formula explains dependent as y and independent value of x . The regression curve is defined by b and a intercept it where ε is the error.

2.9 Artificial Neural Network (ANN)

Artificial neural networks (ANNs), which learn from a training set, are used to predict drought. Multilayer perceptron (MLP) models with one input layer, a hidden layer, and an output layer are often employed. The backpropagation algorithm is used to train the MLP models, which adjusts the connectivity weights using error convergence approaches to provide the desired output for a given input. In this process, the mistake at the output layer in the backpropagation algorithm model propagates backward to the input layer via the network's hidden layer to produce the intended result (Chowdhury, 2017; Ferreira et al., 2019; Morid et al., 2007).

To mimic drought, numerous ANN models are tested with varied weather factors such as average temperatures and rainfall. The result for all models is the Standardized Precipitation-Evapotranspiration Index (SPEI) values. A correlation matrix between distinct input variables and SPEI is generated to examine the relative importance of different input variables on SPEI results. Seventy percent of the data is utilized to train the models, with the remaining 30% used for testing after enhancing the number of nodes in the hidden layer and the network topology. During training, the networks start with random weights and then perform a one-pass backpropagation method at each time step, which includes a forward pass that

propagates the input vector through the network layer by layer and a backward pass that updates the gradient descent rule's weights.

Backpropagation has some drawbacks, such as becoming stuck in local minima and sluggish convergence, yet it is still the most used supervised learning paradigm for ANNs. ANNs can predict drought with reasonable accuracy when utilizing this methodology, making it a viable tool for drought prediction (Dastorani & Afkhami, 2011; Rumelhart et al., 1986).

2.10 Random Forest (RF)

(R.F.) introduced the RF modR.F., which uses a technique called “bagging” to ensemble a collection of decision trees with controlled variance. The RF modR.F. is frequently used in regression and forecasting applications and employs random binary trees trained using a subset of the training dataset via the bootstrapping technique. The RF modR.F. employs two views to train multiple decision trees: the sample dimension and the feature dimension. Combining the voting outcomes of multiple decision trees, the RF modR.F. mitigates the issue of decision trees being prone to over-fitting. Similar to the Bagging model, the core algorithm in the RF modR.F. is trained in parallel. The RF modR.F. is helpful in modeling hydrological features and provides high-precision forecasting without overfitting the data set. When using the RF model as a tree-based ensemble technique, it is essential to evaluate the significance of climatic factors (Wang et al., 2019). For more information on the RF ensemble algorithm, please refer to (Islam et al., 2017; Lotfirad et al., 2022; Park et al., 2018)

2.11 Support Vector machine (SVM)

The Support Vector Machine was initially developed for classification, but it has been extended to regression problems using support vector regression (SVR). This approach uses kernels to minimize the distance to training data and limit model complexity, resulting in better generalization capacity for data not in the training set. In SVR, the SVM regression-based model has two steps. First, the kernel function transforms independent variables such as remote sensing drought indices and weather or climate indices from the original space to a high-dimensional feature space. Second, a linear model is constructed in the new feature space to minimize errors. The most widely used kernel function in the literature is the Radial Basis Function (RBF), which has two hyperparameters: the penalty factor C and the kernel width γ . Both are tuned using a grid search method. This SVM regression-based model for drought prediction offers high generalization capacity and reduces the risk of

overfitting to training data. Further information about the SVM algorithm can be found in (Vapnik, 1995).

3. Results

3.1 Temporal drought trend of SPI and SPEI

MMK trend tests have been utilized in Rangpur Meteorological Station from 1979 to 2020 at the 95% confidence interval to visualize drought trends. The devastating potential trend of Z statistics is found for $SPI_{24} > -8.40$. However, another periodical trend is also found in extreme measurement including $SPI_{12} > -5.60$, $SPI_{09} > -4.50$, $SPI_{06} > -2.75$ and $SPI_{06} > -2.70$, respectively (Figure 2). It must be acknowledged that the MMK trend outcome for SPI is in a heavy upward direction from short-term seasonal to annual variation.

On the other hand, the MMK test for all periodic observations of SPEI reveals a rise in negative trends (Figure 2). A severe drought trend is deemed to control of $SPEI_{24}$, which is $Z > -11.20$. In addition, other time-framed trend suggest the gradual lower Z statistics compared with $SPEI_{24}$. However, $SPEI_{12}$ is the second highest potential drought trend, with z statistics of > -8 . Another periodical trend is also found in relatively decreased trend statistics, indicating the 09 months drought trend $Z > -6.80$, the 06 months drought trend $Z > -5.10$, 03-month period drought trend $Z > -4.80$ (Figure 2). Therefore, the primary comparison between SPEI and SPI trend consequences depicted Z statistics is more intense for SPEI where SPI fabricated relatively reduced trend statistics.

3.2 Relationship between SPI and SPEI

Monthly SPI and SPEI were estimated at five-time scales (3, 6, 9, 12 and 24 months) for the selected station from 1979 to 2020 (Figure 3). The average SPI and SPEI were then calculated and estimated for Pearson correlation coefficient at the 95% confidence interval to characterize drought episodes in the Rangpur district. All SPI and SPEI indices are shown to be significantly associated with an upward trend. It demonstrates a strong relationship ranging from 95% to 98%, where SPI and SPEI 03 score lowest of 95%. However, an increased heavy correlation is found for 9 months, and 12 months among all other indices, including $r = 0.98$. Future correlation estimation suggests that it will increase drought severity for both SPI and SPEI across all indices. Consequently, all future drought indicators exhibited a possible rising trend comparable to the present, with a range of r^2 0.95 and 0.97. Only the 3-month period scored a lower severity of $r^2 = 0.90$ compared to all other future

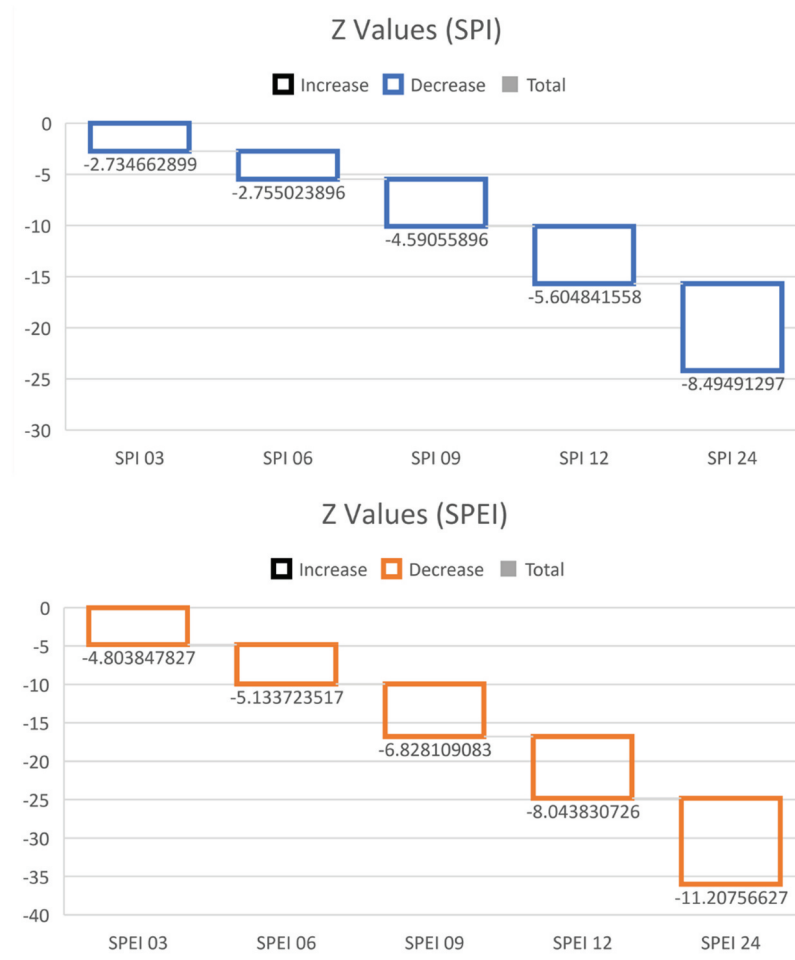


Figure 2. Trend values of Z statistics for SPI and SPEI. Source: Author's own calculation

relationship indices. Thus, the overall scenario indicates a strong association between SPI and SPEI for both present and future variability. Moreover, the 3-month drought-relationship is vulnerable compared to all other indicators, namely 6,9,12, and 24-month periods.

3.3 Comparative assessment between SPI and SPEI

A comparative evaluation was conducted using simple linear regression. The output of the study includes (Figure 4). SPI was unable to identify severe droughts in the Rangpur region based on the brief droughts that occurred in 1983, 1987, 1993, 2002, and 2012; yet, it was able to identify extended periodic drought episodes (in 1984, 1993, 1995, 2007, 2011, and 2013). At this location, this is obvious as SPEI is much more equipped than SPI to recognize instances of drought. It is crucial to understand that SPI exaggerated in many years, including 1994, 1995, 2002, 2007, and 2012. In the years 1989, 1993, 1994, 1995, 1996, 1998, 2004, 2009, 2011, 2014, 2015, and 2016, the SPI could not distinguish between harsh long-term droughts and mild droughts. There were extreme droughts in 1989,

1993, 1994, 1995, 1996, 1998, 2004, 2009, 2011, 2014, 2015, and 2016. These years were identified as severe drought years by SPEI but not by SPI.

3.4 Drought prediction

In this study, three different models were utilized for prediction purposes, namely Random Forest (RF), Support Vector Machine (SVM), and Artificial Neural Networks (ANN). Model accuracy evaluation is a crucial phase in machine learning model development as it enables one to assess the model's performance in terms of prediction. Model performance is evaluated using various metrics such as Mean Squared Error (MSE), Mean Absolute Error (MAE), Root Mean Squared Error (RMSE), and Coefficient of Determination (R-Squared). The training model was developed using 70% of the data, and the testing model was built using the remaining 30%. As we have measured in our previous section, SPEI provides the most accurate drought results, so we have utilized SPEI values for drought prediction. Thus, testing outcomes are also evaluated to demonstrate the prediction results.

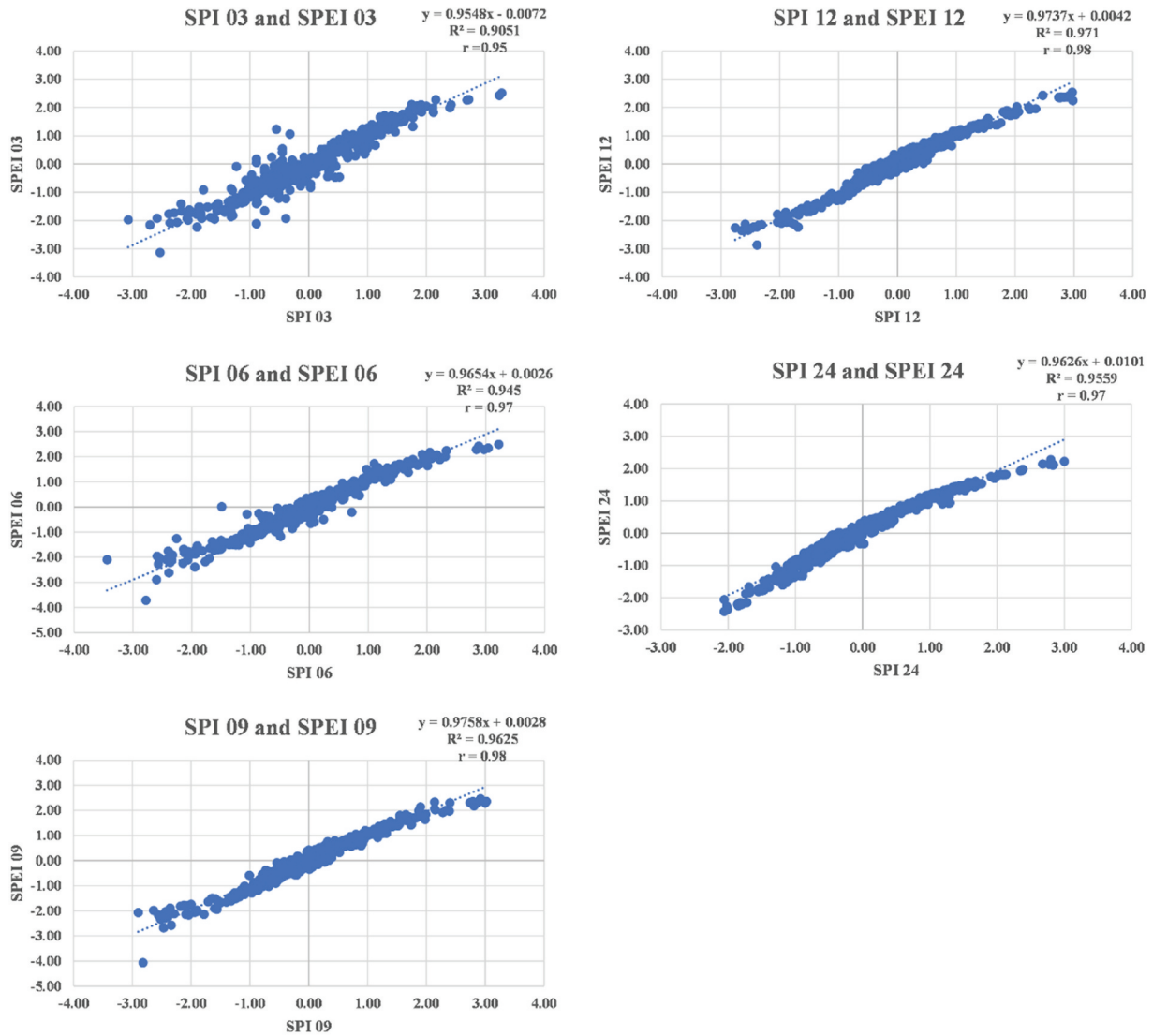


Figure 3. Correlation between SPI and SPEI for the 3-, 6-, 9-, 12- and 24-months. Source: Author's calculation.

In some cases, the R-Squared value may be harmful, indicating that the model is not well-fitted with the datasets. However, for this study evaluation, a negative R-Squared value is considered as 0 statistics. Therefore, we have found 0 statistics for all the model of long-term drought prediction, namely SPEI 09, SPEI 12, and SPEI 24.

Additionally, to visualize 0 statistics clearly, we coloured it bold in blue. However, for short-term drought prediction of SPEI 03 only ANN provides a significant value of $0.06 r^2$ where considerable RMSE, MSE and MAE is found. Thus, for the seasonal drought prediction ANN found superior over RF where RMSE, MSE and MAE is found lower in ANN model. This ANN result also indicates the increase of seasonal drought of 24%. All models, however, demonstrated low prediction accuracy (negative R2 values). Overall, the models' performance varied according to the SPEI value, and more study is required to increase the precision of drought prediction models.

4. Discussion

The primary objective was to evaluate drought trends, determine the relationship between SPI and SPEI, and develop predictions. Drought was either steady, diminished, or exacerbated in various ways, according to the outcomes. Therefore, SPI and SPEI exhibited a strong positive correlation, and both indices rose (Figure 3). Similar to this study (Adnan et al., 2018), determined that SPEI has adequate capacity to assess drought conditions in Pakistan. These results are congruent with those of (Homdee et al., 2016). In contrast to this study (Mahmoudi et al., 2019), recognized SPI and effective drought index (EDI) as Iran's primary and second choice dryness detection indicators (Kamruzzaman et al., 2019). Kamruzzan discovered that EDI significantly improved than SPI for detecting longer and shorter periodic shortages in Bangladesh. This anomaly may be attributed to our project region's plain, low-lying geography, which promotes seafloor airflow inconsistencies. Concerning intensity and amplitude, the association across SPI and SPEI were

highest for 3-, 6-, 9-, 12-, and 24-month timelines (0.95, 0.97, 0.98, 0.98, and 0.97, correspondingly) (Figure 3). Following the period of transition, severe droughts became increasingly prevalent. This could lend support to (Cook et al., 2014) assertions that the greenhouse effect has exacerbated drought in various parts of the world in the twenty-first century.

SPI failed to detect severe drought occurrences in this region in 1987, 1989, 2002, 2004, and 2009 (Figure 4). SPI and SPEI have done an excellent performance in detecting short-term droughts in the middle section of the country. According to Malik (Deye et al., 2016), a severe and widespread drought afflicted most of Southeast Asia between 1999 and 2002, coinciding with this study's findings. In contrast, SPI failed to recognize protracted severe dryness in 1993, 1995, 2014, and 2016, misclassifying these years as moderate water shortages (Figure 4). There were significant water shortages in 1989, 1993, 1994, 1995, 1996, 1998, 2004, 2009, 2011, 2014, 2015, and 2016 (Figure 4). SPEI proved that it can identify these periods as severe drought, although SPI did not. This research highlighted droughts during the 1980s, 1990s, 2000s, and 2010s. These findings can be linked to those of (Miah et al., 2017) and (Rahman & Lateh, 2016a). An exaggeration or underestimating could be due to rainfall being the primary entry parameter in the SPI computation. In contrast, SPEI fared exceptionally well in detecting previous droughts in Rangpur. According to (Stagge et al., 2015; Tan et al., 2015), SPEI exhibited actual drought conditions since higher air temperatures or little precipitation led in significant evaporation. Thus, SPEI can reveal the effects of rains and evaporate on dryness (Vicente-Serrano et al., 2014); PET showed excellent agreement with the meteorological drought indicator (Joetzjer et al., 2013). This discovery validates significant findings from (Miah et al., 2017). According to the researchers, PET is a major component of the hydrology in the western part of Bangladesh and also in Rangpur, and thus its variation is mostly driven by precipitation (Alamgir et al., 2015; Islam et al., 2017)

Table 3 and supplementary figures (s1, and s2) depict visual examination and appraisal of prediction systems in the context of conventional output indicators as regression analysis. In addition, only the ANN prediction model performed well for short-term and seasonal drought prediction of SPEI 03 and SPEI 06, which is an upward trend. Such results are consistent with those of (Ali et al., 2017; Poornima & Pushpalatha, 2019). In the Asian region, ANN provides better result for drought prediction than SVM and ANN (Mokhtar et al., 2021; Park et al., 2018). The Table 2 indicates drought index based on SPEI values.

In contrast, several studies suggested that future changes in precipitation could reduce the severity of drought in the respected research region (Bari

et al., 2016; Habiba et al., 2013b). But the predictive power of each model was subpar. While some models performed slightly better in some scenarios than others, overall, all models performed below expectations. More research is required to increase the accuracy of drought prediction models, especially for longer term droughts like SPEI 12 and SPEI 24. However, despite being a minor outlier, most environmental warming forecasts imply that temperatures will climb significantly in Bangladesh, and the dry periods' dryness (November to May) intensify in the foreseeable future (Hasda et al., 2020; Rahman & Lateh, 2016b; Yaseen et al., 2021). Bangladesh makes drought forecasting challenging due to its nonlinear climate pattern. When considering the relevance of regression analysis, the drought prediction model varies according to the periodic and yearly phases. The drought prediction accuracy of the created systems was examined in the existing kinds of the literature and quantitative parameters in the region (Bari et al., 2016; Mahsin et al., 2012) and concluded to be superior (Alamgir et al., 2020; Hasda et al., 2020; Kamruzzaman et al., 2019; Rahman et al., 2021). Drought has a considerable negative effect on different species and subsistence agriculture; for instance, a severe drought in 1978 decreased rice productivity by 2 million tons (Habiba et al., 2013a). During 1982, dryness impacts were much and more significant than twice flooding effects (Umma & Rajib, 2014).

Additionally, more than 2 million hectares of farmland were lost to the droughts of 1997 (Habiba et al., 2013b). This area's drought caused crop loss or decreased output, leading to the acquisition of many farm products (PR, 1995). Since rice is grown in Bangladesh throughout the summertime, the country may face water scarcity, mostly in times of inclusive growth, because of the waste of precious groundwater (Mainuddin et al., 2020). However, in recent years, an exceptional drought event has developed in Rangpur district, primarily affecting the northern zone's northern section. Drought incidence has been particularly severe during SPI 24 and SPEI 24, SPI 12 and SPEI 12, SPI 06 and SPEI 06. Rangpur district is in conventional measures of extreme drought for 6 months, 12 months, and 24 months, but 3 months and 9 months droughts are also becoming more severe.

Furthermore, according to the trend study, and prediction the Rangpur district is at critical drought risk. Similar findings have been acknowledged by (Abdullah, 2014; Rahman et al., 2021). The reason for these changes is that precipitation varies spatially, temporally and seasonally (Rahman & Lateh, 2016a). Therefore, monsoon season wind and precipitation, as well as Himalayas features, play a role in the dry spell

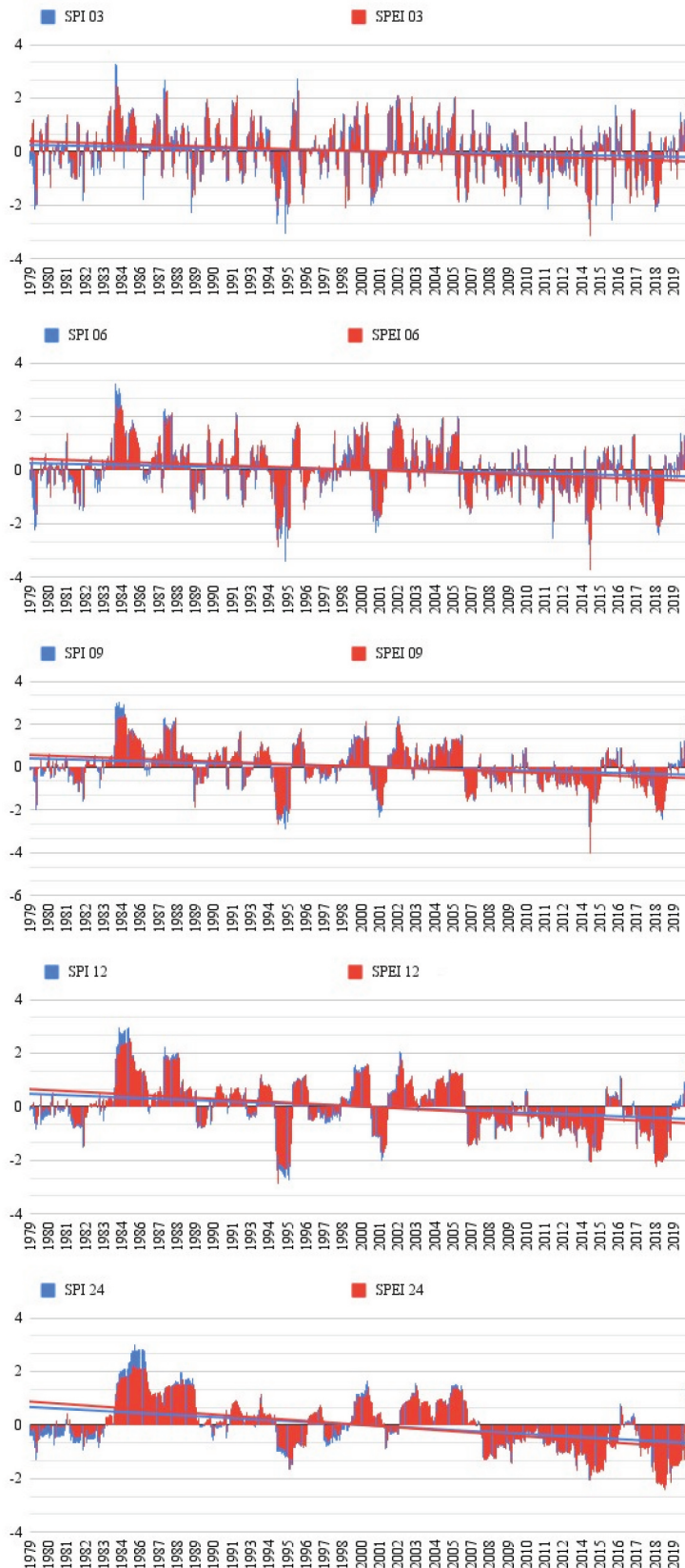


Figure 4. Comparative assessment between SPI and SPEI. Source: Author's calculation

and precipitation patterns of Bangladesh and Rangpur (Das, 2021). El Nino also significantly affected Bangladesh's climate, resulting in decreased precipitation and drought in different seasons (Islam et al., 2021). Another very probable proof is that greenhouse

gas emissions are responsible for Rangpur's persistent dryness (Mora et al., 2018). However, it should be noted that extreme dryness persists because of a shortage of precipitation and the prevalence of water stress regions in Rangpur (Rahman & Azim,

Table 3. The prediction outcome of drought.

SPEI		SVM (Training)	SVM (Testing)	RF (Training)	RF (Testing)	ANN (Training)	ANN (Testing)
SPEI 03	RMSE	0.87968	0.94284	1.0385	1.0481	0.85976	0.87125
	R2	0.23	-0.10	-0.08	-0.36	0.26	0.06
	MSE	0.77383	0.88895	1.0786	1.0985	0.73918	0.75907
	MAE	0.70715	0.73716	0.83896	0.85388	0.6836	0.68255
SPEI 06	RMSE	0.93909	0.96639	0.9927	1.1563	0.91918	0.97256
	R2	0.12	-0.23	-0.76	0.02	0.16	0.24
	MSE	0.88189	0.93392	0.98545	1.337	0.8449	0.94587
	MAE	0.73435	0.74689	0.80116	0.90589	0.72274	0.76796
SPEI 09	RMSE	0.94458	1.0672	1.0727	1.1617	0.96318	1.0683
	R2	0.09	-0.75	-0.17	-1.07	0.06	-0.75
	MSE	0.89223	1.1388	1.1506	1.3495	0.92771	1.1412
	MAE	0.73814	0.86251	0.83782	0.92703	0.75958	0.86711
SPEI 12	RMSE	0.95255	1.1277	1.0567	1.2821	0.9687	1.1015
	R2	0.04	-1.27	-0.19	-1.94	0.00	-1.17
	MSE	0.90734	1.2717	1.1166	1.6438	0.93837	1.2134
	MAE	0.732	0.95535	0.81839	1.049	0.75617	0.94519
SPEI 24	RMSE	0.8114	1.3762	0.99313	1.632	0.87045	1.4436
	R2	-0.01	-3.06	-0.51	-4.71	-0.16	-3.47
	MSE	0.65837	1.894	0.98631	2.6633	0.75768	2.0841
	MAE	0.65787	1.2343	0.81963	1.4055	0.70871	1.2998

Source: Authors own calculation.

2022; Rahman et al., 2021). The correlation between the amount of heat in the water and the amount of rainfall is a possible reason for the prolonged extreme drought (BS, 2007). Bangladesh's northwestern region, including the Rangpur district, endures drought during dry periods because of the Farakka Barrage, which hinders the Padma river's movement. This causes 57% severe water scarcity (WBB, 1998). Consequently, temperature and rainfall would probably rise in the coming years, possibly resulting in a water stress deficiency in the Rangpur area district (Chowdhury et al., 2020; Rahman & Azim, 2022). Conversely, rising temperatures, on the other hand, might reduce annual precipitation in the coming decades, which could lead to extended dry spells in Bangladesh (Rahman & Lateh, 2016a).

Nevertheless, drought monitoring and prediction is essential for crop management, water scarcity management, water demand and supply evaluation, city and regional planning. The limitation of the study includes only one station-based drought characterization and prediction. However, the study area contains only one station it can be further studied based on remote sensing approaches. Therefore, more specifically further study may analyse more aspects and properties of drought in Bangladesh to find the most effective drought mitigation strategy on the reported trends with public perception.

5. Conclusion

This investigation examined the temporal tendencies of drought based on the relationship and comparative evaluation between SPEI and SPI. Also, predictions of temporal drought occurrence in Rangpur district during 1979–2020 have been performed. Therefore, several techniques have been

used to identify drought trend exploration, relationship and comparative assessment evaluation as well as prediction. The statistical measurement of linear regression was used as a part of parametric test to detect the trend of monthly rainfall and temperature before applying non-parametric test including Modified Mann-Kendall for trend analysis for acquiring SPEI and SPI or drought as well as dry and wet period variability. As a part of correlation, Pearson correlation coefficient have been utilized. Additionally, ANN, RF and SVM machine learning method was employed to predict future variability of drought trend for SPEI. Significant downward negative trend is found for SPI 03, 06, 09, 12 and 24 which corresponds to $Z > -2.70$, $Z > -2.75$, $Z > -4.50$, $Z > 5.60$, $Z > -8.40$, respectively. Additionally, SPEI has an extensive adverse growth of negative trend compared with SPI Z statistics that includes SPEI 03 > -4.80 , SPEI 06 > -5.13 , SPEI 09 > -6.82 , SPEI 12 > -8.04 and SPEI 24 > -11.20 . A significant correlation is observed for all the indices of SPEI and SPI indicates $r = 0.95$, $r = 0.97$, $r = 0.98$, $r = 0.98$ and $r = 0.97$ for 03, 06, 09, 12 and 24-months period, respectively. Comparative assessment distinguished SPEI is extensively suitable for drought variability measurement with the comparison of SPI. Moreover, no predictive model performed well for annual drought estimation. Hence, seasonal drought will increase of 24% for SPEI 06. The results stress the significance of improving control of hydro resources and reallocating crop production policy frameworks in Bangladesh as a means of dealing with disasters. This long-range prediction is designed to aid policy in arranging for more effective drought monitoring, rainwater collection, and cultivation practices.

Disclosure statement

No potential conflict of interest was reported by the author(s).

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